

Modelling Ancient Earth Climate: *Methods & Models*

Prof. Alan M. Haywood

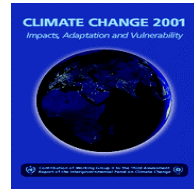
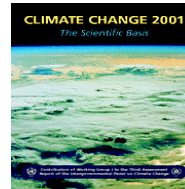
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with contributions from Paul Valdes, Ulrich Salzmann, Victoria Peck, Steve Hunter & Jane Francis



Why?

- Understand the dynamics of warm climates
- Test Earth System Models



Primary Research Focus in Climate Change Science

- Simulation of the historical or near-historical record
- Analysis of the observed record of variability
- Projection for the next 100 years

Greatest Strengths

Spatial and temporal character of the Observations.
Measurement of physical quantities that define the state of the atmosphere and ocean.

Greatest Weaknesses

Sense of change.
Sense of the integration of the Earth System.

In contrast: A Research Focus in Earth History

Greatest Strengths

Spectacular sense of change (*Furry Alligator Syndrome*)
True integrated system response

Greatest Weaknesses

Proxies rather than state variables
Limited spatial and temporal resolution

“The greatest weaknesses in a research focus on the modern record are the greatest strengths of Earth System History”

We Should Worry

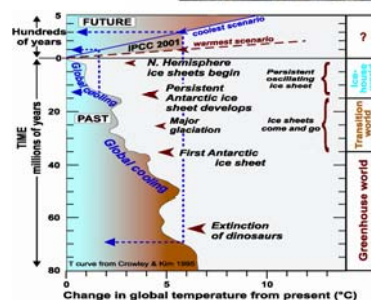
IPCC Climate Sensitivity: Roughly 1.5 to 4.5 C globally averaged surface temperature increase for a doubling of carbon dioxide.

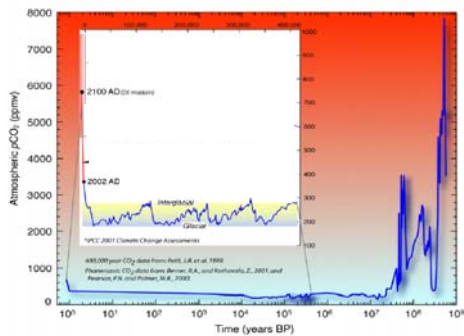
Hundreds of GCM experiments have been completed for time periods throughout the Phanerozoic using a wide variety of climate models.

Many experiments focused on either glacial climates or warm climates (the extremes).

“There is no legitimate example of a climate model simulation in which the past climate conditions were overestimated”

Climate History





Atmospheric carbon dioxide through the Phanerozoic and projected into the next century. Inset modified from figure constructed by S. Porter (2003). L. Sohl constructed the Phanerozoic portion (2003).

1.1 Introduction: What exactly is a "model" ?

model n. [Fr. Modele, It. Modello, from L. modellus]
A miniature representation (small measure) of a thing, with the several parts in due proportion.

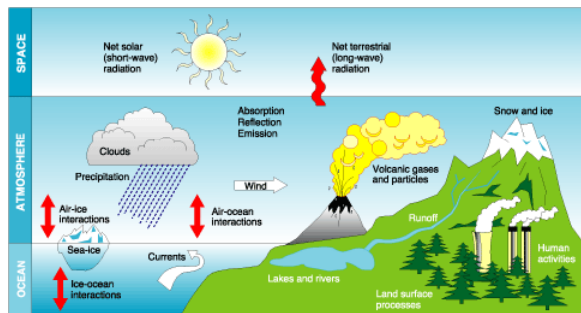
- A model is only a "representation" of reality (e.g. a street plan of reality)
- Good modellers know the **strong** AND **weak** points of their models
- "Modelling" (English) and "Modeling" (American)
- Some quotations:

"All models are wrong, but some are useful" – George Box

"The purpose of models is not to fit the data but to sharpen the questions" – Samuel Karlin

"A theory has only the alternative of being right or wrong. A model has a third possibility, it may be right, but irrelevant." – Manfred Eigen

The Climate System



General Circulation Models

Model needs to simulate albedo, emissivity and general circulation.

Use "first principles"

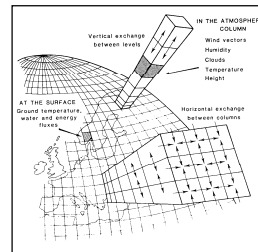
Newton's Laws of Motion

1st Law of Thermodynamics

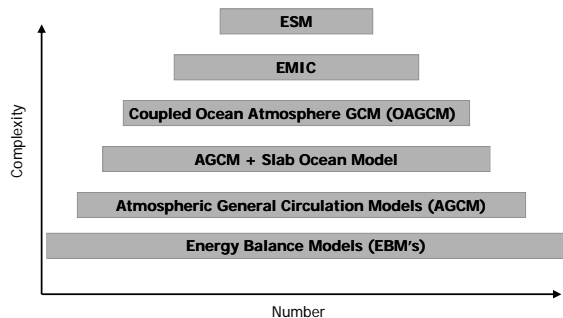
Conservation of Mass and Moisture

Hydrostatic Balance

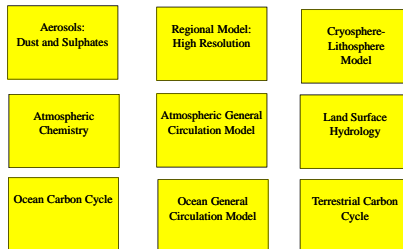
Ideal Gas Law



Spectrum of Climate Models

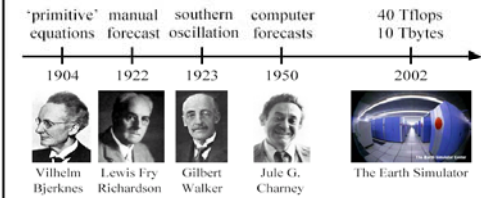


Components of an Earth System Model



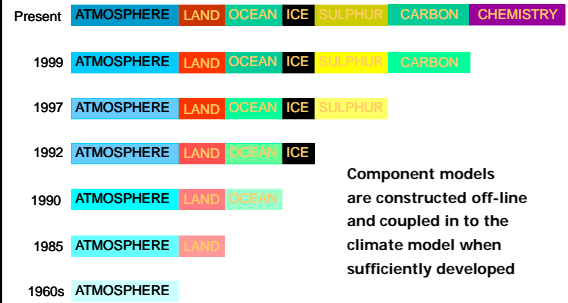
Brief history of numerical modelling

1.2 Brief history



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Development of Met. Office Climate Models



Physical basis of climate models

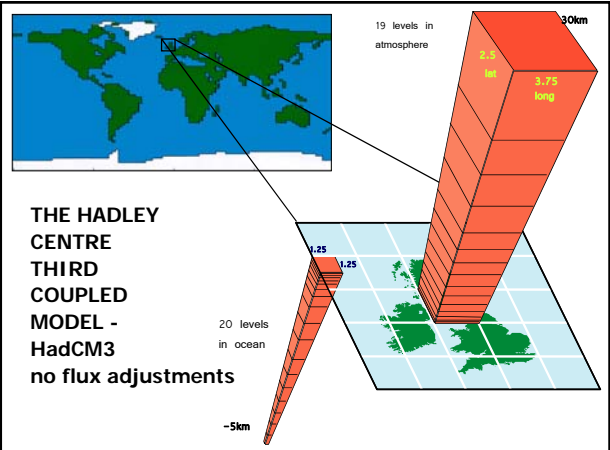
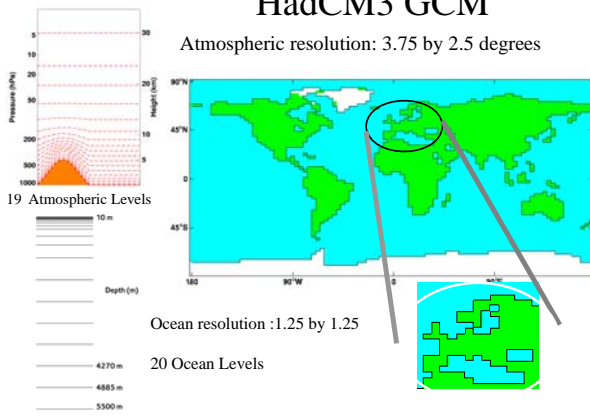
- The atmosphere is a fluid on a rotating planet:
 - Drag at the surface and within the atmosphere affects the momentum budget
 - Water vapour evaporates from the surface, condenses to form clouds and heats the atmosphere when it is lost through precipitation
 - Heating from solar radiation and cooling from thermal radiation
- Models therefore need to include equations for:
 - 3 components of wind (or vorticity & divergence), including Coriolis and drag
 - equation of state and conservation of water
 - thermodynamics, including heating by condensation and radiation

Physical basis of climate models

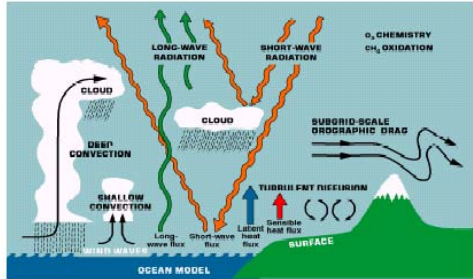
- The ocean is also a fluid, but incompressible. It is heated by solar radiation and cooled by evaporation and thermal emission from the surface. No internal heating, but salinity strongly affects the density and hence the circulation
- Additional models have been developed to include the land surface, cryosphere, atmospheric chemistry and aerosols, carbon cycle etc
- **Processes that are sub-grid in scale are modelled by parametrizations**

HadCM3 GCM

Atmospheric resolution: 3.75 by 2.5 degrees



Parametrized processes in the ECMWF model

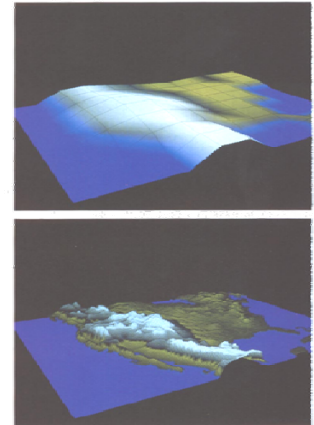


Representation of orography; the importance of resolution

The upper figure shows the surface orography over North America at a resolution of 480km, as in a low resolution climate model.

The lower figure shows the same field at a resolution of 60km, as in a weather forecasting model.

Remember that orographic processes are highly non-linear.

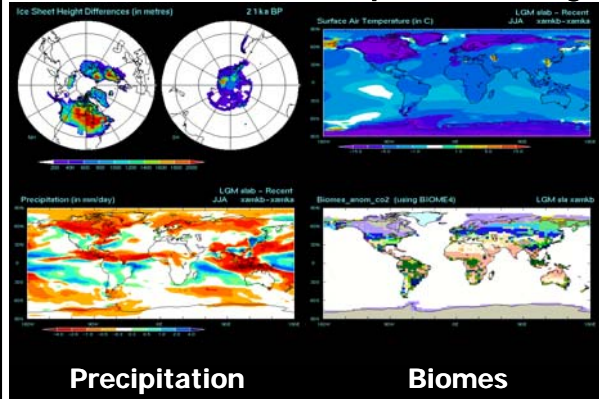


So.....

The horizontal and vertical resolutions of climate models need to be high enough to avoid numerical errors and to resolve the basic dynamical and transport processes

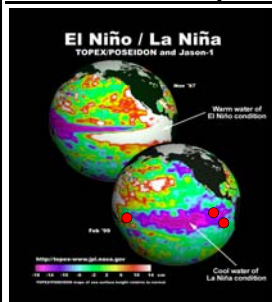
There is a trade-off between resolution and computing time, but model resolutions are increasing continually, as more computer power becomes available

Ice Sheet Elevation Temperature Change



Case Studies

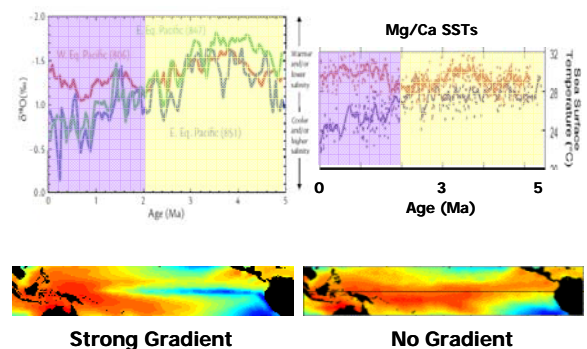
1. Palaeo ENSO (El Niño Southern Oscillation)

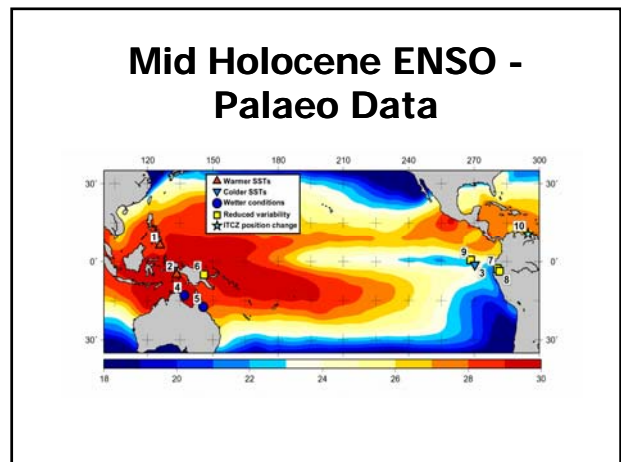
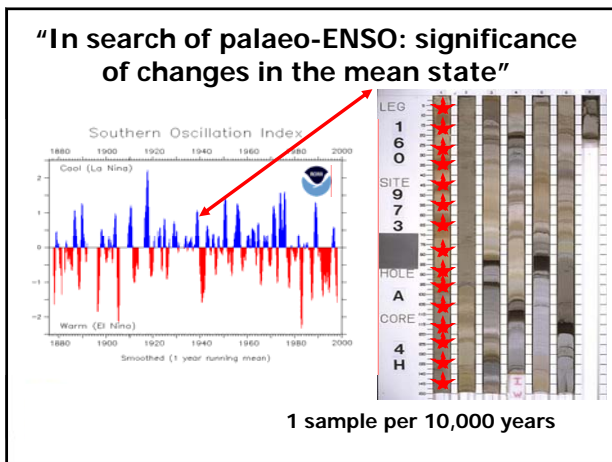
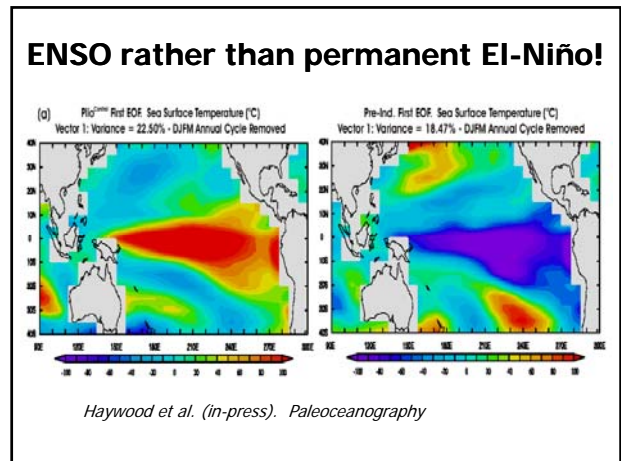
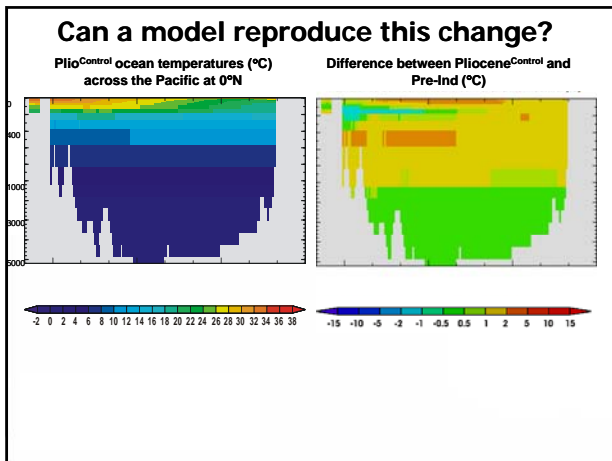


- Coupled ocean-atmosphere phenomena
- Involves large scale fluctuations in a number of oceanic/atmospheric variables (e.g. sea surface temps. & sea level pressure)
- El Niño & La Niña opposite extremes of ENSO

The Pliocene: a Permanent El Niño-like state?

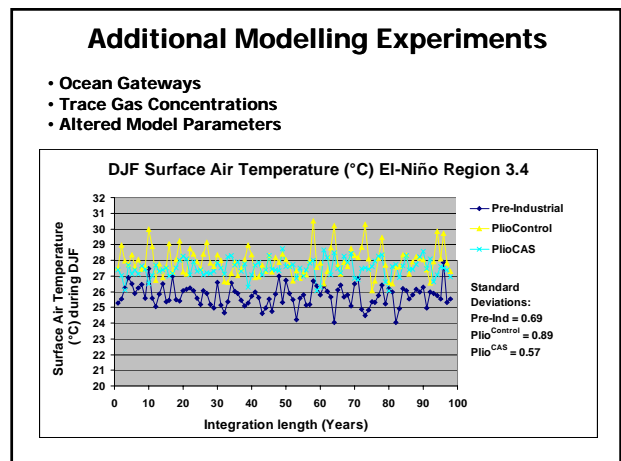
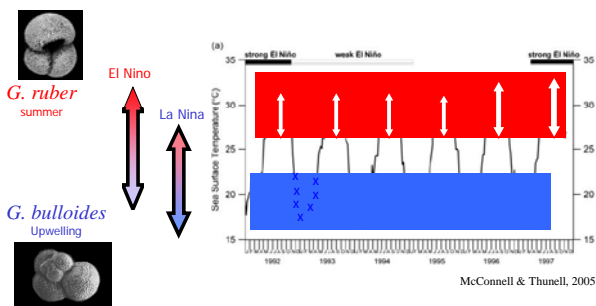
(Wara et al., 2005; Philander & Federov, 2003)

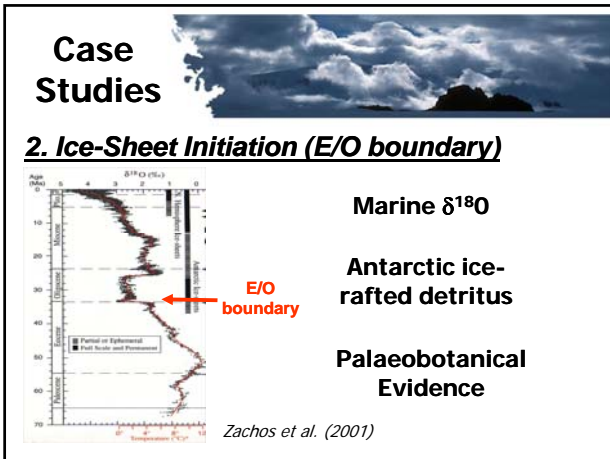




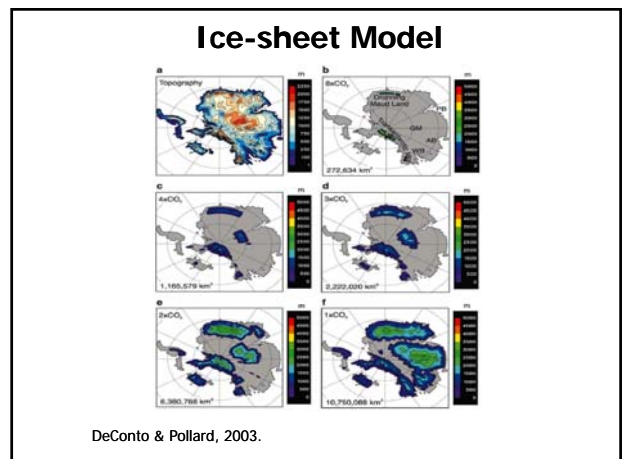
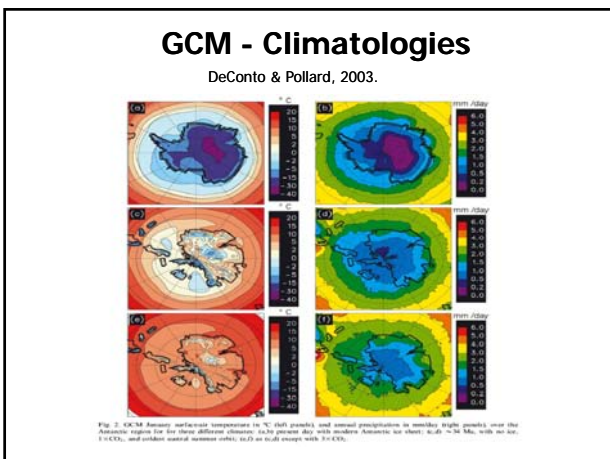
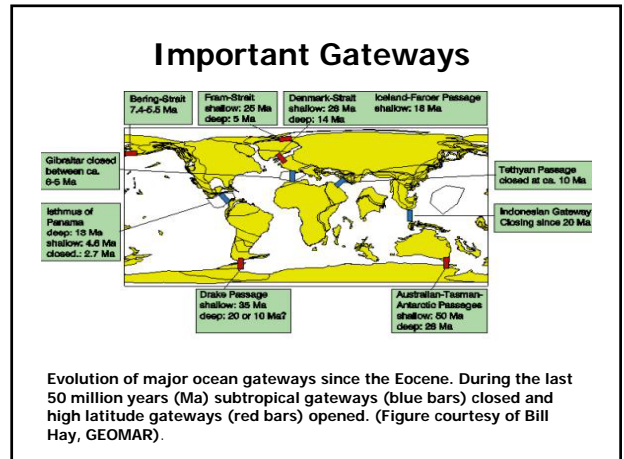
ENSO - Interannual variability

Single specimen analyses however, may provide insight to the range of seasonal extremes within time slices, similar to modern studies in the Gulf of California.

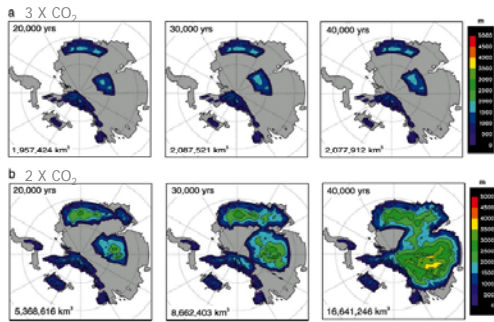




- ### Causal Mechanisms for Antarctic Ice-Sheet Initiation
- Ocean Gateways
 - Antarctic elevation
 - Declining atmospheric CO₂ concentrations

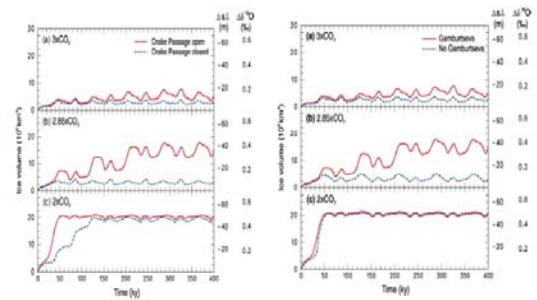


Ice-sheet Model



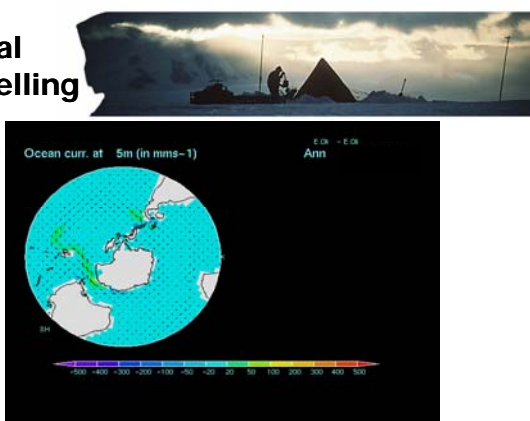
DeConto & Pollard, 2003.

Sensitivity Tests



DeConto & Pollard, 2003.

Initial Modelling



Case Studies

3. Cretaceous Climates & Ice-Sheets



Palaeobotanical Evidence

Cretaceous forest 120 million years ago on the Antarctic Peninsula. reconstruction based on PhD of Jodie Howe, University of Leeds/BAS, painted by Robert Nichols.

Evidence for large, rapid sea-level changes (Miller et al., 2005)

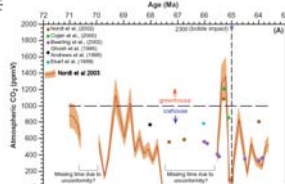


- Evidence for eustatic nature
- Pace and magnitude suggest glacial origin.
- Suggest moderate-sized ice sheets (5 - 10 × 10⁶ km³).
- Paced by Milankovitch forcing.

CO₂ levels through the Maastrichtian

- 2 greenhouse episodes 1000-1400 ppm
- But suggestions of CO₂ low-points at ~70 and 66 Ma.

(Nordt et al., 2003; Beerling et al., 2002).



How to create a Maastrichtian model

Change solar output ~0.6% less than present

CO₂ (and other gases) 4 x pre-industrial (but could be 2x to 8x).

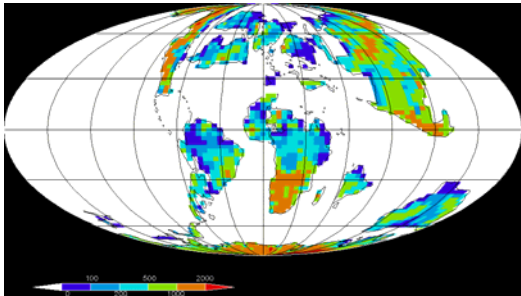
Volcanic activity Assume same as today.

Change in orbit Same as present, but perform sensitivity simulations

Palaeogeography Including sea-level/ orography/ bathymetry/land ice

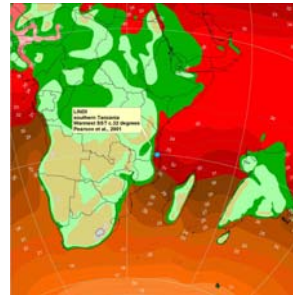
Previous modelling also required prescription of vegetation, and sea surface temperatures (or ocean heat transport) but this is no longer needed.

Maastrichtian Orography



At climate model resolution. Original palaeogeographies from Paul Markwick

Coupled Ocean-Atmosphere Simulation: Comparison to Oxygen Isotopes

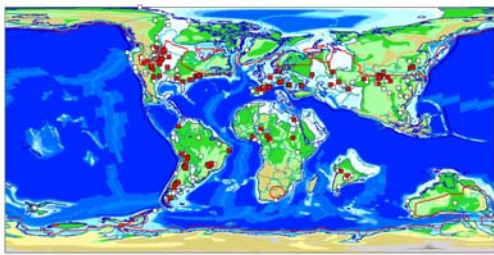


Model predicted temperatures approx. 10C at 1000m, 8C at 2000m, and 7C at 3000m

c.f. temperatures from 14C to 7C from D'Hondt & Arthur (2002)

Paul Pearson's Maastrichtian data

Coupled Ocean-Atmosphere Simulation: Comparison to Vertebrate Data



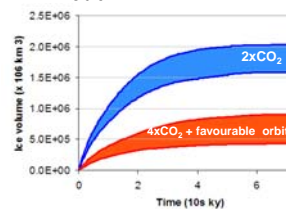
Red squares= all crocs, Orange= Dinosaurs, White = Other Vertebrates

Model predicted cold month mean shown by 5C contour (red) and 0C (blue)

Paul Markwick's database

Ice Sheets in a Greenhouse World

- Suite of HadCM3 derived palaeoclimates
 - 2, 4, and 6 x CO₂
 - Further runs being carried out including 1 x CO₂
- Comparison against climate proxy database
- Climate then used to drive a BAS ice-sheet model.



2xCO₂ - ~ 2 x 10⁶ km³ ice

The Future: Data & Models Combine

Limiting Factors:

- Available computer power
- Model sophistication (resolution etc.)
- Small community ("every department needs a pet modeller")
- Language barriers & a lack of communication
- A new generation of scientists to act as the interface

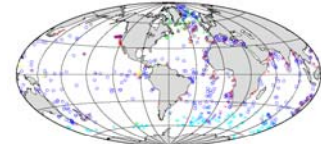
A couple of examples what modellers need from the geological community

- Ocean temperatures
- Land cover

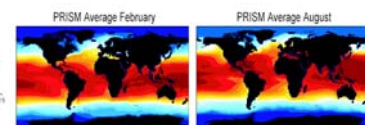
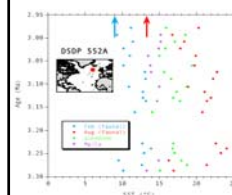
"Deep-time perspectives on climate change: marrying the signal from computer models and biological proxies (Eds. M. Williams, A. M. Haywood, J. Gregory and D. Schmidt) The Micropalaeontological Society & the Geological Society of London"

Oceans:

Quantitative temperature estimates derived from multi-proxy studies PRISM & MARGO

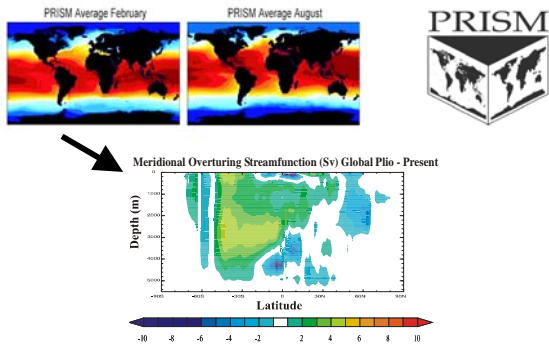


Kucera et al. (2005) QSR Vol. 24, 813-819



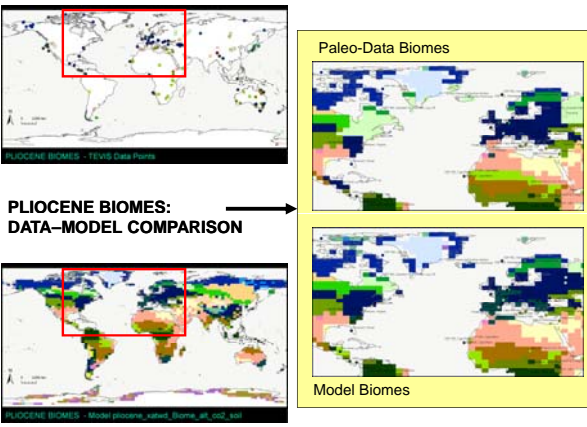
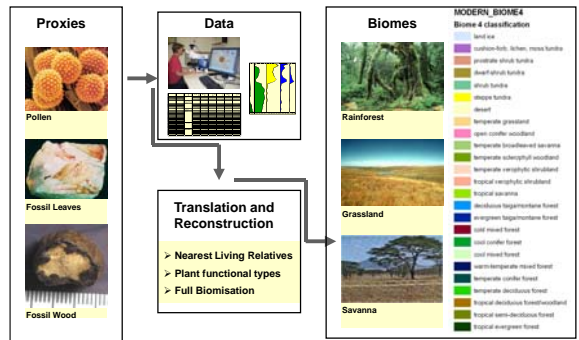
H. Dowsett (in-press).

New 3-D Ocean Temperature Reconstructions: multiple equilibrium?



Land Cover:

Vegetation and Biomes derived from paleobotanical studies



Summary

- Why Model?
- Need for Integration
- How its Done
- Permanent El-Niño?
- Importance of ACC on EAIS kick off?
- Glaciation in a Greenhouse word? **NO PROBLEM**

